Extreme Gammon 2 Activation Key

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The annual and 3-month average June-August SST anomalies over the contiguous United States show a spatial pattern similar to that of heavy precipitation (Figure 2.6, bottom). Averaged over US land areas, the 21st century SST increase of 0.38°C will likely bring additional frequency and intensity of extreme heat events over the United States. This pattern is attributed to the combination of increased greenhouse gas and aerosol forcing and the resulting higher levels of water vapor in the atmosphere. While greenhouse gas-driven warming is primarily responsible for raising SSTs at the surface, aerosol-driven warming provides the primary contribution to the increased levels of water vapor in the atmosphere, which in turn is the primary factor in boosting heavy precipitation frequency and intensity. As a result of these changes, an increase in the frequency of some extreme weather and climate related events and processes is projected for the U.S. and other countries. 2 One consistent theme across all assessments is that more frequent extreme events pose increasing risks to human health, infrastructure and environments. Advancing understanding of the link between human-induced climate change and changes in extreme events is an important priority. The center of the extreme temperature time series reflects the overall distribution of extreme temperatures, which in turn depends on the frequency distribution of daily mean temperatures in monthly mean temperature series (Stephens et al. 2003; Eq. (3.22) in Sect. 3.1). However, the tail-end of the time series represents extreme events and is related to the weather extremes of low or high daily maximum temperatures that occur infrequently. These extreme events are not only the rarest extreme events but are also the most damaging for crops and human activities. The tail-end of the time series of maximum temperature is highly sensitive to climate change, which can alter the probability of extreme events during specific seasons. Lower extreme maximum temperatures, for example, increase the frequency of days that are near the upper tail of the time series of maximum temperatures during the summer months, which are more likely to result in more extreme maximum temperature events (e.g., Stephens et al. 2003; Stephens et al. 2007; Stephan et al. 2009; Box 3.1, Fig. 4.1). Other factors, such as the duration of the heat wave, may also contribute to more extreme maximum temperature events (Donnelly et al. 2012). The assessment of these changes in future extreme maximum temperature events should take into consideration the response of the tail-end of the time series to recent and projected climate changes (Donnelly et al. 2012; Lukas et al. 2017; Box 3.3).

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The annual and 3-month average June-August SST anomalies over the contiguous United States show a spatial pattern similar to that of heavy precipitation (Figure 2.6, bottom). Averaged over US land areas, the 21st century SST increase of 0.38°C will likely bring additional frequency and intensity of extreme heat events over the United States. This pattern is attributed to the combination of increased greenhouse gas and aerosol forcing and the resulting higher levels of water vapor in the atmosphere. While greenhouse gas-driven warming is primarily responsible for raising SSTs at the surface, aerosol-driven warming provides the primary contribution to the increased levels of water vapor in the atmosphere, which in turn is the primary factor in boosting heavy precipitation frequency and intensity. As a result of these changes, an increase in the frequency of some extreme weather and climate related events and processes is projected for the U.S. and other countries. 2 One consistent theme across all assessments is that more frequent extreme events pose increasing risks to human health, infrastructure and environments. Advancing understanding of the link between human-induced climate change and changes in extreme events is an important priority. The center of the extreme temperature time series reflects the overall distribution of extreme temperatures, which in turn depends on the frequency distribution of daily mean temperatures in monthly mean temperature series (Stephens et al. 2003; Eq. (3.22) in Sect. 3.1). However, the tail-end of the time series represents extreme events and is related to the weather extremes of low or high daily maximum temperatures that occur infrequently. These extreme events are not only the rarest extreme events but are also the most damaging for crops and human activities. The tail-end of the time series of maximum temperature is highly sensitive to climate change, which can alter the probability of extreme events during specific seasons. Lower extreme maximum temperatures, for example, increase the frequency of days that are near the upper tail of the time series of maximum temperatures during the summer months, which are more likely to result in more extreme maximum temperature events (e.g., Stephens et al. 2003; Stephens et al. 2007; Stephan et al. 2009; Box 3.1, Fig. 4.1). Other factors, such as the duration of the heat wave, may also contribute to more extreme maximum temperature events (Donnelly et al. 2012). The assessment of these changes in future extreme maximum temperature events should take into consideration the response of the tail-end of the time series to recent and projected climate changes (Donnelly et al. 2012; Lukas et al. 2017; Box 3.3). 5ec8ef588b

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